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Local air pollution from oil rig emissions observed during the airborne DACCIWA campaign

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Abstract. In the framework of the European DACCIWA (Dynamics–Aerosol–Chemistry–Cloud Interactions in West Africa) project, the airborne study APSOWA (Atmospheric Pollution from Shipping and Oil platforms of West Africa) was conducted in July 2016 to study oil rig emissions off the Gulf of Guinea. Two flights in the marine boundary layer were focused on the floating production storage and offloading (FPSO) vessel operating off the coast of Ghana. Those flights present simultaneous sudden increases in NO₂ and aerosol concentrations. Unlike what can be found in flaring emission inventories, no increase in SO₂ was detected, and an increase in CO is observed only during one of the two flights. Using FLEXPART (FLEXible PARTicle dispersion model) simulations with a regional NO₂ satellite flaring inventory in forward-trajectory mode, our study reproduces the timing of the aircraft NO₂ enhancements. Several sensitivity tests on the flux and the injection height are also performed, leading to the conclusion that a lower NO₂ flux helps in better reproducing the measurements and that the modification of the injection height does not impact the results of the simulations significantly.

1 Introduction

Crude oil extraction from offshore platforms brings raw gas mixed with oil to the surface. Gas flaring is used to dispose of this natural gas in cases where the infrastructure to export it does not exist. This process emits a mixture of trace gases like carbon dioxide (CO₂), carbon monoxide (CO), sulfur dioxide (SO₂), nitrogen oxides (NO_x) and particulate matter. Its impacts concern both the ecosystems (Nwaugo et al., 2006; Nwankwo and Ogagarue, 2011) and the air quality (Osuji and Avwiri, 2005). The pollutants can be transported into the free troposphere (Fawole et al., 2016b) or can reach coastal cities in the marine boundary layer (MBL). Flaring emissions can be derived from remote-sensing techniques (Elvidge et al., 2013) by the widely used Visible Infrared Imaging Radiometer Suite (VIIRS) Nightfire satellite product (Deetz and Vogel, 2017). Oil and gas extraction activities are, however, uncertain in terms of emitted quantities (Tucella et al., 2017), and direct measurements are necessary.

In Africa, most of the studies focused on the environmental impact of Nigerian oil platform emissions (e.g. Anomohanran, 2012; Hassan and Kouhy, 2013; Asuoha and Osu Charles, 2015; Fawole et al., 2016a), as it is one of the five countries with the highest flaring amounts (Elvidge et al., 2015). Other countries in the Gulf of Guinea are nevertheless affected by environmental problems. The DACCIWA (Dynamics–Aerosol–Chemistry–Cloud Interactions in West

Africa; Knippertz et al., 2015) project conducted fieldwork in southern West Africa (SWA) in 2016 to investigate the impact of anthropogenic emissions, notably on air quality. One part of the project included airborne measurements of flaring emissions in order to fill the data gap on oil extraction activities in SWA. Those in situ measurements are, as far as we know, the first reported data on oil extraction activities in this region.

We use here the Lagrangian transport model FLEXPART (FLEXible PARTicle dispersion model; Stohl et al., 2005) in combination with an inventory of emissions dedicated to flaring emissions and created in the framework of the DACCIWA project (Deetz and Vogel, 2017) to reproduce the aircraft measurements. We focus on evaluating the sensitivity of the model to two parameters: the emission flux and the injection height of the flaring emissions. After a brief description of the DACCIWA project and the platform in Sect. 2, we present in Sect. 3 the model and the flaring emission inventory we used as well as the estimation of the flaring plume injection height. We discuss the modelling results in Sect. 4.

2 The DACCIWA–APSOWA project

2.1 Description of the campaign

The EU-funded project DACCIWA focuses on the coupling between dynamics, aerosols, chemistry and clouds (Flamant et al., 2017). The research campaign was set up in June–July 2016 and undertook activities ranging from airborne measurements to running atmospheric numerical models. A map showing the location of the research site is presented in the Supplement (Fig. S1).

Our Atmospheric Pollution from Shipping and Oil Platforms of West Africa (APSOWA) project complementing DACCIWA seeks to characterise gaseous and particulate pollutants emitted by oil platforms off the coast of the Gulf of Guinea by dedicated flights conducted with the DLR (Deutsches Zentrum für Luft- und Raumfahrt) research aircraft Falcon 20. Different instruments for gas and particulate measurements were deployed onboard. We focus on CO, NO₂, SO₂ and aerosol measurements during two flights on 10 and 14 July. Both CO and NO₂ were measured by SPIRIT (SPectromètre InfraRouge In situ Toute altitude; Catoire et al., 2017). This infrared absorption spectrometer uses continuous-wave distributed-feedback room-temperature quantum cascade lasers (QCLs) allowing on-line scanning of mid-infrared rotational–vibrational lines with spectral resolution of 10^{-3} cm^{-1} . In the present campaign, the ambient air is sampled in a multipass cell with a path length of 134.22 m and in micro-windows around the 2179.772 cm^{-1} line for CO and the 1630.326 cm^{-1} line for NO₂. A home-made software using the HITRAN 2012 database (Rothman et al., 2013) is used to deduce the total

molecule abundance. The overall uncertainties are 4 ppbv for CO and $0.5 \text{ ppbv} + 5 \%$ for NO₂ at 1.6 s time resolution.

SO₂ measurements are performed using a pulsed fluorescence SO₂ analyser (Thermo Electron, Model 43C Trace Level). Ultraviolet (UV) light is absorbed by SO₂ molecules in the sample gas which become excited and subsequently decay to a lower energy state. The emitted light is detected by a photomultiplier tube and is proportional to the SO₂ concentration in the sample gas. The time resolution of the measurements is 1 s, with a moving average of 30 s to smooth the data. The lower detection limit is 0.315 ppbv. The instrument was multipoint calibrated before and after the campaign.

Total aerosol concentrations were measured with a butanol-based condensation particle counter (CPC; CPC TSI 3010, modified for aircraft use; Schröder and Ström, 1997; Fiebig et al., 2005). The particle counter was mounted inside the fuselage behind the Falcon isokinetic aerosol inlet. The large particle cut-off diameter imposed by the inlet has been found to be between 1.5 and $3 \mu\text{m}$, depending on flight altitude (Fiebig, 2001). The lower cut-off diameter of the CPC was at $\sim 10 \text{ nm}$. Particle losses due to diffusion effects were minimised by using a standard 5.75 L min^{-1} bypass flow to the instrument. Using the particle loss calculator described in von der Weiden et al. (2009), we estimate the particle losses in the tubing to be less than 10 % for the relevant size range. Flight sequences inside clouds are known to cause sampling artefacts and have therefore been removed from the dataset.

2.2 Flight planning and FPSO description

The first floating production storage and offloading (FPSO) vessels were installed in Indonesia in 1974. Since this year, their number has steadily increased (Shimamura, 2002). The concept of those platforms based on ship structure makes the development of small-size oil fields and the exploitation of them further from the coast and thus in deeper water possible (Shimamura, 2002). Those platforms have other advantages (Shimamura, 2002): they are faster to build than other floating structures, they have inbuilt storage capability and thus do not necessitate pipelines, and they are movable and easily implantable on another oil field. Because of all those reasons, the FPSO systems will continue to develop. Among the techniques used to dispose of the gas associated with the crude oil extraction, one is the flaring, consisting of burning the gas in an open flame through a stack. This leads to a mixture of emitted gases which can reach the free troposphere if the meteorological conditions are favourable (Fawole et al., 2016b). The two targeted flights, on 10 and 14 July, consisted of about 3–4 h of meandering transects through emission plumes in the MBL about 300 m a.s.l.) off the coast of West Africa, from Côte d'Ivoire to Togo. We focus on the flights in the vicinity of the FPSO Kwame Nkrumah platform, on the Jubilee Field, off the coast of Ghana ($4^{\circ}35'47.04 \text{ N}$, $2^{\circ}53'21.11 \text{ W}$). It measures 330 m in

length by 65 m in width, and its height up to the top of the chimney is estimated to be 112 m a.s.l. During the campaign, no contact with this FPSO could be obtained in order to have more information on its functioning.

3 Method

3.1 Lagrangian particle dispersion modelling

The Lagrangian model FLEXPART (Stohl et al., 2005) is used to study the transport of the emitted plume from the FPSO in the MBL. It simulates long-range transport and dispersion of atmospheric tracers released over time by computing trajectories of a large number of tracer particles. Model calculations are based on meteorological data from the European Centre for Medium-Range Weather Forecasts (ECMWF), ERA-INTERIM L137 (Dee et al., 2011), extracted every 3 h and with a horizontal grid mesh size of $0.5^\circ \times 0.5^\circ$. The calculations are performed in forward-dispersion mode with the model version 9.0. The particles are released with the chemical properties of NO_2 , CO and SO_2 using constant emissions from the Deetz and Vogel (2017) inventory during 7 h, with a spin-up of 5 h, allowing the model to be balanced independently from the initial conditions. During the simulation, the NO_2 - and SO_2 -like particle mass is lost by wet and dry deposition and by OH reaction (concentrations from GEOS-CHEM model; technical note FLEXPART v8.2, <http://flexpart.eu/downloads/26>, last access: November 2018), which allows a lifetime of about 3 h at 298 K in the MBL for NO_2 . CO-like particle mass is only lost by OH reaction.

3.2 Flaring emission inventory

Our study is based on a new gas-flaring emission inventory developed for the DACCWA project (Deetz and Vogel, 2017). This inventory provides emissions of CO, CO_2 , NO, NO_2 and SO_2 for June–July 2014 and 2015, and we use a 2016 updated version for the period of the APSOWA–DACCWA campaign. It is based on remote-sensing observations using VIIRS nighttime radiant heat in combination with combustion equations from Ismail and Umukoro (2016). The emission estimation method is described in detail in Deetz and Vogel (2017). Only the assumptions of interest for our study are indicated: first, the natural gas composition includes 0.03 % H_2S (Sonibare and Akeredolu, 2004). Second, this composition measured in the Niger Delta is valid for West Africa in general. Third, the source temperature is deduced from the VIIRS measurements on a monthly mean. For the FPSO platform, it is set to 1600 K, which is a good order of magnitude, since the flame temperature can be as high as 2000 K (Fawole et al., 2016b). Fourth, for such a temperature, NO_2 is considered to be a primary pollutant coming from the rapid conversion of NO close to the source, and the inventory does not include any later transformation

of the species. The CO, SO_2 and NO_2 fluxes estimated with this method for our 2 d of interest for the FPSO platform are 1.1×10^{-1} , 4.23×10^{-5} and $7 \times 10^{-2} \text{ kg s}^{-1}$, respectively. Note that the highest uncertainties (+33 %; –79 %) associated to the estimation of gas-flaring emissions arise from the parameters required in the combustion equations, e.g. the gas composition, the source temperature and the flare characteristics.

3.3 Estimation of the flaring plume injection height

The oil flaring emissions are generally emitted at higher temperatures than the temperature of the surrounding environment, which implies an important role for the buoyancy at the stack exit. Indeed, the buoyancy corresponds to a density ratio between the air parcel and its colder surrounding environment and leads to the rise of this parcel under the influence of gravity. This effect is to be distinguished from the momentum effect, defined as the product of an element mass by its velocity, which can be neglected for such a high-temperature plume (Briggs, 1965). The buoyancy raises the plume above its initial injection height and can lead to a source height that can be several times higher than the real height of the stack (Arya, 1999). The calculation required to determine the rise of a plume depends on the wind conditions and the atmospheric stability. Before determining those parameters, we determine the MBL height that defines the part of the troposphere in which we flew. The MBL is about 680 m a.s.l. on 10 July and 582 m a.s.l. on 14 July, according to the European (ECMWF) and US (NCEP) operational forecasts. The wind conditions are determined by the Falcon 20 measurement system. An average of the wind speed measurements is calculated during the flight period in the vicinity of the platform (see Sect. 4.1). For 10 July, an 8 min mean wind speed of $9.4 \pm 0.5 \text{ m s}^{-1}$ was calculated, where the standard deviation represents the natural variability, which is larger than the measurement uncertainty. For 14 July, a 7 min wind speed mean of about $6.6 \pm 0.7 \text{ m s}^{-1}$ was calculated. Thus, for both days the conditions are windy.

Concerning the potential temperature for 10 and 14 July, the atmosphere is considered to be stable with a positive potential temperature gradient. With the previous parameters defined, it is possible to calculate the plume injection height ΔH (in m) by using Eq. (1), reported by Briggs (1965, 1984):

$$\Delta H = 2.6 \left(\frac{F_b}{us} \right)^{1/3}, \quad (1)$$

with F_b as the buoyancy effect (in $\text{m}^4 \text{ s}^{-3}$) defined in Eq. (2), u as the wind speed (in m s^{-1}) and s as the stability parameter

Table 1. Flux and injection height for the reference control run (CTRL) and for the sensitivity tests (STs) for each day of flight.

Run name	Date of flight	NO ₂ flux (kg s ⁻¹)	SO ₂ flux (kg s ⁻¹)	CO flux (kg s ⁻¹)	Injection height (m)
CTRL	10 Jul 2016	0.07	4.23×10^{-5}	0.11	27
	14 Jul 2016				
ST1	10 Jul 2016	0.07	Not included	Not included	68
	14 Jul 2016				77
ST2	10 Jul 2016	0.035–0.05	Not included	Not included	27
	14 Jul 2016				
ST3	10 Jul 2016	0.035–0.05	Not included	Not included	68
	14 Jul 2016				77

(in s⁻²) defined in Eq. (3) (Briggs, 1965):

$$F_b = g \frac{\Delta T}{T_s} w r^2, \quad (2)$$

$$s = \frac{g}{T} \left(\frac{\partial T}{\partial z} \right) + \Gamma. \quad (3)$$

T , T_s , w and r are the absolute temperature of the ambient air, the absolute temperature of the stack gases, the vertical velocity of the effluent at the stack exit and the radius of the stack, respectively. ΔT is the difference between the two temperatures. In Eq. (3), g is the gravitational acceleration, z the altitude and Γ the adiabatic lapse rate. Both the temperature at the stack exit and the effluent velocity are taken from Deetz and Vogel (2017). The calculated plume injection height ΔH is 27 m for 10 and 14 July. Briggs' algorithm underestimates the plume rise in the stable atmospheric condition (Akingunola et al., 2018). Therefore, another method has been tested to determine the injection height that is based on VDI 3782 (1985). For a stable atmosphere, the injection height ΔH (in m) of the plume is determined by Eq. (4):

$$\Delta H = 74.4 \times Q^{0.333} \times u^{-0.333}, \quad (4)$$

with u as the wind speed (in m s⁻¹) and the unit of the coefficient (74.4; in s⁴ kg⁻¹ m⁻²). The heat flow Q (in units of MW) (Eq. 5) is defined in Deetz and Vogel (2017) as

$$Q = H \frac{1}{f}, \quad (5)$$

with H as the radiant heat observed by VIIRS (in units of MW) and f as the fraction of radiated heat set to 0.27 by Deetz and Vogel (2017) after having averaged the f values given in Guigard et al. (2000). The calculated plume injection heights are, according to Eq. (4), 68 and 77 m for 10 and 14 July, respectively.

The two methods of injection height calculation do not give similar results. The control run (CTRL; Table 1) will use Briggs' algorithm with an injection height of the particles of 27 m for both days. A sensitivity test (ST1; Table 1)

will be performed to see the impact on the results of the injection height in the model by using the injection height from VDI 3782 (1985).

4 Results and discussion

4.1 Description of the measurements

Figure 1 presents the part of the flights in the vicinity of the FPSO platform. During the flight on 10 July, the flaring plume was crossed several times (Fig. 1a). It led to four NO₂ peaks between 12:33 and 12:45 UTC at an altitude of around 300 m (Fig. 1c). Four simultaneous peaks of aerosols were measured (Fig. 1c). Neither simultaneous CO peaks (within its lower detection limit, 0.3 ppbv; Catoire et al., 2017) nor SO₂ peaks (within its lower detection limit, 0.3 ppbv; see Sect. 2.1) were detected in the plume of the FPSO platform for this flight. No peak was measured during a second series of transect at a higher altitude (around 600 m; Fig. 1a).

During the flight on 14 July, three NO₂ peaks were measured during the first transects between 10:48 and 11:00 UTC at around 300 m altitude downwind of the FPSO vessel (Fig. 1b). Those peaks were simultaneously measured with aerosol and CO peaks (Fig. 1d), but still no SO₂ peak was detected. Knowing that SO₂ comes from H₂S combustion (Sonibare and Akeredolu, 2004), these results suggest that a gas composition of 0.03 % of H₂S induces an emission of SO₂ concentration lower than the detection limit of the instrument from 3 km of our measurements or that the natural gas composition given by Deetz and Vogel (2017) for the Niger Delta is different from that in Ghana for those two flights. The presence (or lack) of CO peaks is discussed in Sect. 4.3.

Moreover, SPIRIT allows measurements every 1.6 s. Considering the case on 10 July where we have the maximum aircraft speed (118 m s⁻¹) and the shortest peak (lasting about 16 s), the plume width is larger than the measurement e -fold time multiplied by the flight speed. Thus, for all the narrow

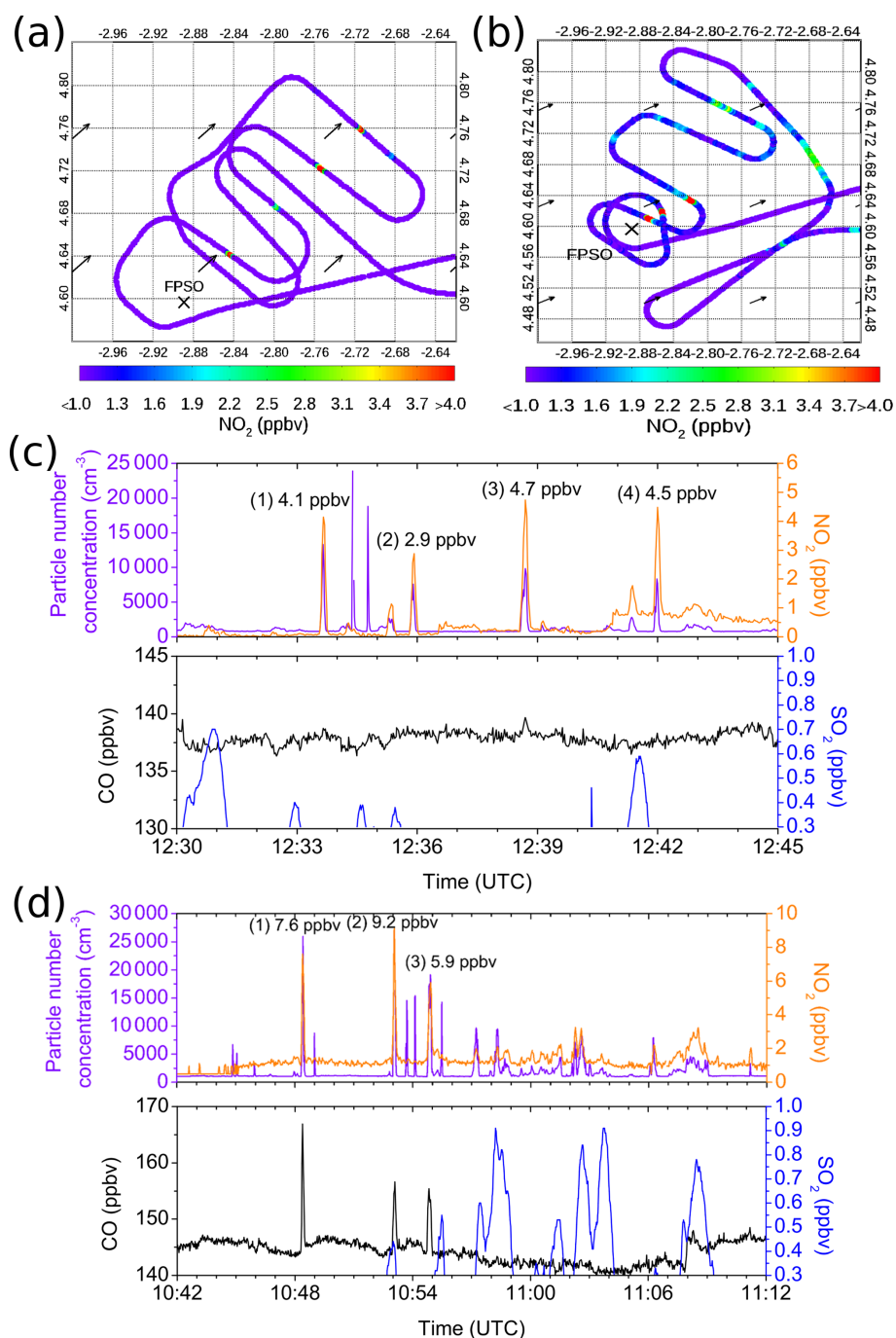


Figure 1. (a) NO₂ concentration as a function of the flight trajectory downwind of the FPSO plume for 10 July. The black arrows show the wind direction (from ECMWF). (c) NO₂, aerosol, CO and SO₂ concentrations as a function of time, restricted in a part of the flight trajectory in (a). The peaks studied are labelled by a number (from 1 to 4). (b) and (d) are similar to (a) and (c) for 14 July.

peaks, the maximum plume concentration is real, not a plume diluted with its surrounding environment.

4.2 FLEXPART simulations of the flaring emissions

In order to confirm that the peaks detected by the aircraft instruments correspond to the flaring emissions from the platform and to simulate them, forward trajectories are calculated using FLEXPART. First, as a reference, a simulation

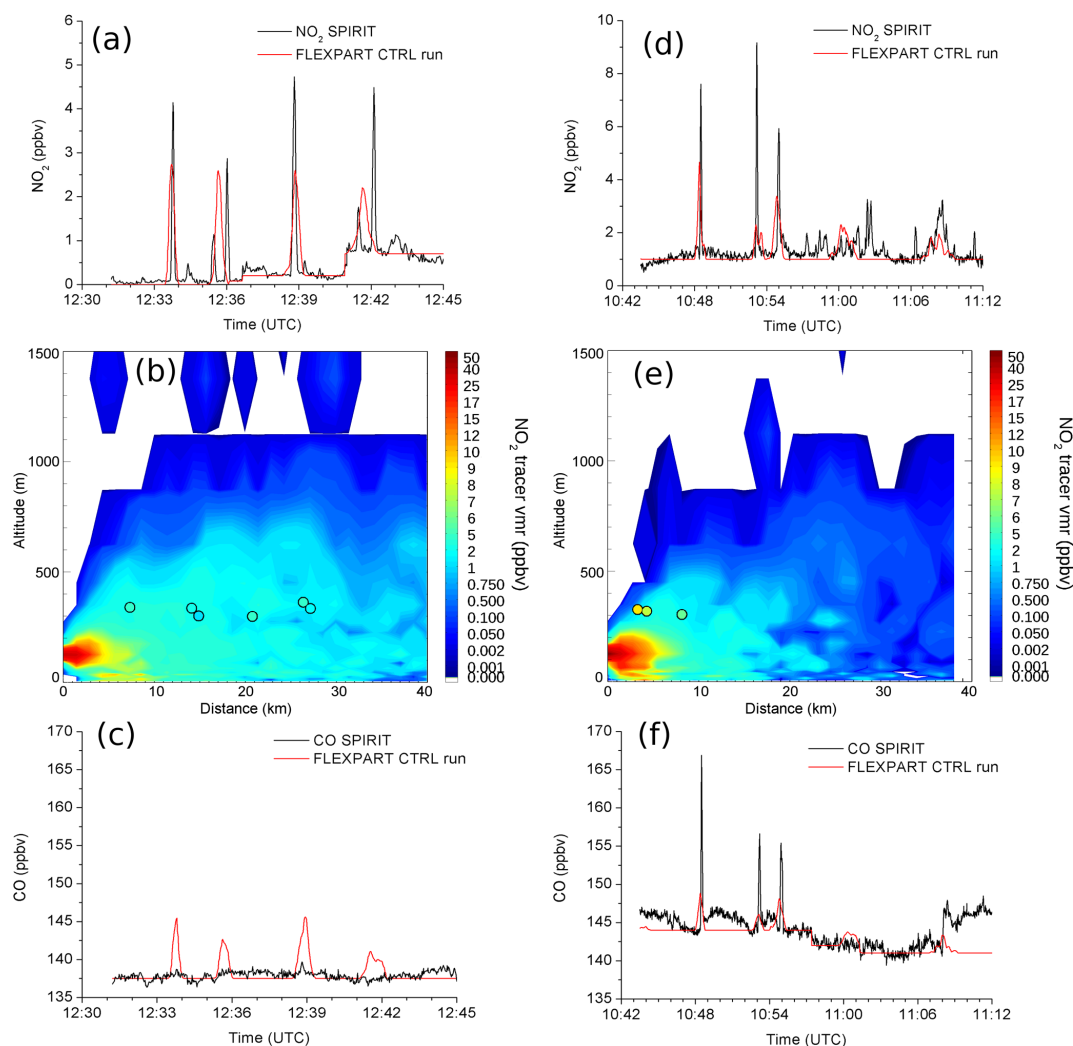


Figure 2. Left column: 10 July. (a) NO_2 concentration as a function of time for SPIRIT measurements (black) and FLEXPART CTRL run simulation (red). (b) Vertical section of the simulated NO_2 concentration (in vmr – volume mixing ratio with units of ppbv) in the plume (CTRL run) as a function of distance from the source and altitude. The coloured circles correspond to the measurement peaks. (c) CO concentration as a function of time for measurements (black) and FLEXPART CTRL run simulation (red). Right column: 14 July. (d), (e) and (f) similar to (a), (b) and (c).

called CTRL is run in which the NO_2 flux for the FPSO platform is taken from Deetz and Vogel (2017), i.e. $7 \times 10^{-2} \text{ kg s}^{-1}$, and the injection height is 27 m, calculated with Briggs' method (see Sect. 3.3). A comparison of the wind speed and direction between ECMWF simulations and the Falcon measurements is presented in the Supplement. The wind speed and direction derived from ECMWF agree within 1 m s^{-1} and $\sim 1.8^\circ$, respectively, for 10 July (Fig. S2) and within $\sim 1 \text{ m s}^{-1}$ and $\sim 7^\circ$, respectively, for 14 July (Fig. S3). So the transport is well reproduced in FLEXPART. Figure 2a compares observed and simulated time series of NO_2 concentrations for the flight on 10 July. The four measured peaks are all remarkably well reproduced in time by FLEXPART. Another study by Tuccella et al. (2017), using WRF-Chem, was able to reproduce the simulated peaks si-

multaneously to the measured ones for oil facilities emissions in the Norwegian Sea. A NO_2 background concentration has been added to the last two peaks. This value is an average of the measurements taken outside the plume and is representative of the ambient pollution not taken into account in FLEXPART, as this model only simulates the pollution coming from the platform. Concerning the second and the fourth peak (Fig. 2a), the measurements show two close peaks that FLEXPART cannot simulate individually, leading to a single and broader simulated peak. This is probably due to an error in the dispersion modelling induced by the horizontal and vertical wind field resolution that prevents us from comparing peak-to-peak concentrations. Even with a finer wind field grid mesh of $0.125^\circ \times 0.125^\circ$ (simulation not shown), such close peaks cannot be distinguished, suggesting a still-

insufficient spatial resolution. Instead, the integrated area under each of the measured and simulated plume transects will be compared and presented in Fig. 3, with the percentages representing the relative differences with respect to SPIRIT measurements. Figure 2d also compares the observed and simulated NO_2 concentrations for the flight of 14 July. The simulation gives three peaks concomitant with the three measured peaks of interest.

Comparisons between CTRL run and SPIRIT measurements for both flights (Fig. 3a1 for 10 July; Fig. 3b1 for 14 July) show that the simulated concentrations are always overestimated for all the peaks (percentage larger than 0 %). Sensitivity tests are performed in order to show the FLEXPART response to the flux. Fluxes lower than 0.07 kg s^{-1} are tested from 0.035 to 0.05 kg s^{-1} . Figure 3a1 and b1 show the variations in the differences between the measurements and the FLEXPART simulations with the flux used in FLEXPART for the injection height from Briggs (1965). The plots for both days show a linear response of FLEXPART to the flux increase, with a best estimation reached for 0.04 kg s^{-1} for the flight on 10 July and 0.035 kg s^{-1} for the flight on 14 July. Figure 3a2 and b2 show similar conclusions but for the injection height from VDI 3782 (1985). To determine whether the observed linear relationship between the percentages and the flux or the injection height occurs by chance, a simple F test is performed, assuming that the variances are homogeneous and the results follow a Gaussian distribution. F -statistic coefficients are calculated and compared to the 95 % or 90 % confidence interval with $(1, N - 2)$ (N : total number of results) as degrees of freedom (see values in brackets [u ; $+\infty$] in Fig. 3). If the value F is included in the confidence interval, then the relationship can be considered to be linear. The standard errors on the slope are also added in the plots of Fig. 3.

For the flight on 10 July, the standard error of the slope coefficients and the F test (95 % confidence) show linear relationships between the percentage difference and the flux. Only the results on 14 July for the plot with the injection height of VDI 3782 (1985, panel b2) show a positive F test, but with 90 % confidence. No conclusions can be drawn for the results on 14 July with the injection height of Briggs (1965; Fig. 3b1). In order to show the response of FLEXPART to the injection height, Fig. 3a3 and b3 show the percentages versus the injection height (Briggs, 1965, or VDI 3782, 1985). FLEXPART shows similar results regardless of the injection height used as input whatever the flux used. All the cases show standard error on the slope coefficients larger than the slope itself and an F value not included in the confidence interval (at 95 % confidence, as shown in the figure, and even 90 %, not shown). These results suggest that the differences between the two injection heights are not significant enough with respect to the vertical resolution of the model or that the measurements are too far to be influenced by the changes in this parameter. However, to really conclude about the injection height and to evaluate the flux,

more measurements are needed at different altitudes and distances from the emission source.

Besides the weather conditions and the functioning of the platform, the flight location is also an important parameter to be able to evaluate our measurements. Figure 2b and e show the NO_2 plume simulated with FLEXPART as a function of distance from the source and altitude on 10 and 14 July, respectively. The aircraft measurements, represented by the coloured circles, are located at the upper part of the plume, away from the strongest concentrations. The work carried out is thus limited by the flight trajectories which were too high and too far from the FPSO platform to catch the part of the plume with the highest concentrations. The operational conditions during the flights were complex for the pilots, and safety concerns forced us to respect a minimum flight level (300 m) and a minimum distance from the source. Finally, we found that the NO_2 concentration difference between the measurements and the simulations does not seem to depend on the distance from the source, since the measurements are already too far.

Emissions from oil and gas production vary depending on the operating conditions of the platforms, making them variable and hard to analyse (Law et al., 2017). Considering this, it is possible to have sources of NO_x (and SO_2 if present) from non-flaring combustion processes like the power generation for the facility (Villasenor et al., 2003). No review can be found in the literature to estimate the magnitude of those emissions. As a comparison, during the ACCESS aircraft campaign (Tuccella et al., 2017) in the Arctic, the maximum NO_x mixing ratio associated to oil and gas platforms is about 10 ppbv (Fig. 5 in Tuccella et al., 2017). For the ACCESS campaign, it is known that the facilities were operating under normal conditions, and the flight altitude ranges from 120 to 250 m. The comparison of the FPSO in the Gulf of Guinea with the platforms in Arctic shows that the NO_2 measured during DACCWA campaign is in the range of a normal functioning mode. As mentioned in Zhang et al. (2019), the sources associated to non-flaring combustion are not always negligible, depending on the operating conditions. Thus, we cannot exclude a contribution from another NO_x source overestimating the results, but this contribution is not important when gas-flaring activity is in normal mode (Zhang et al., 2019).

We can use the best simulation on each day to estimate the percentage of pollutants transported inside and above the MBL. In both cases, about 90 % of the pollutants stay inside the MBL and are able to impact the population living along the coastline. Measurements made along the coastline have shown that NO_2 concentrations are generally greater than 2 ppb for suburban sites and greater than 20 ppb near industrial sites (Bahino et al., 2018). Given the wind velocity (from 6.6 to 9.4 m s^{-1}), the air masses attain the coast in 2 to 3 h, which does not allow the transport of significant NO_2 concentrations to impact air quality in this area. This is confirmed when looking at FLEXPART simulations in Fig. 2b

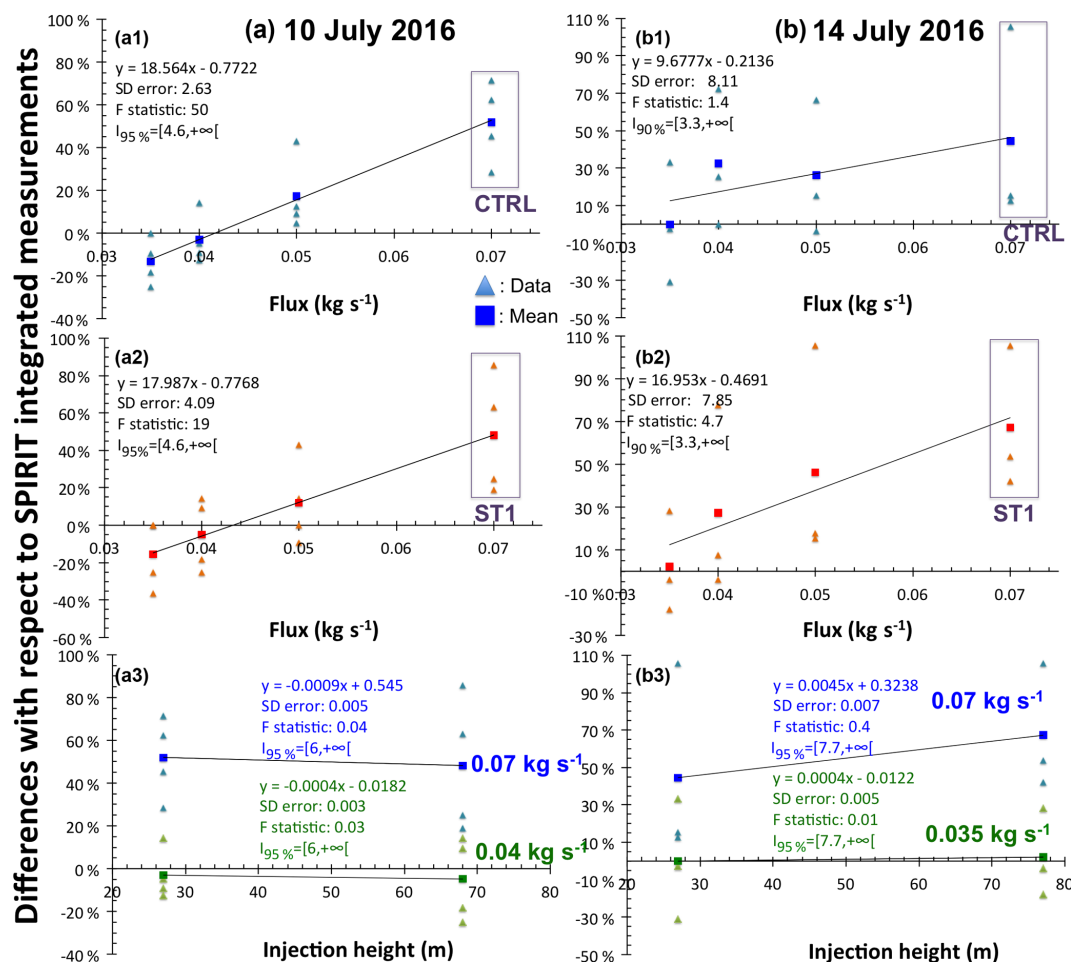


Figure 3. Differences (in %) between FLEXPART simulations and SPIRIT integrated measurements depending on flux or injection height used as input in the model for (a) the flight on 10 July and (b) the flight on 14 July. Panels (a1) and (b1) represent the change in the percentage with the flux by using the injection height from Briggs' algorithm (1965; blue data; i.e. 27 m) and panels (a2) and (b2) with injection height from VDI 3782 (1985; orange data; i.e. 68 m – a2 – or 77 m – b2). Panels (a3) and (b3) represent the change in the percentage with the injection height for the flux from Deetz and Vogel (2017; blue data; 0.07 kg s^{-1}) and for the flux used in the sensitivity tests (green data; 0.04 kg s^{-1} for 10 July in a3 and 0.035 kg s^{-1} for 14 July in b3). For all panels, triangles represent the data for all the peaks measured and squares represent the mean from these data. The slope, standard error values for the slope coefficients, the F statistic and the confidence interval ($I_{95} \%$ or $I_{90} \%$ only for panel b1 and b2) are added for all the plots.

and e. They show that NO_2 concentrations are already very low (< 1 ppbv) from 40 km from the source on 10 July and even closer (from 20 km from the source) on 14 July. With the distance between the coast and the emission source following the wind direction being 70 km, only pollutants with a relatively long lifetime or secondary pollutant like O_3 can impact the air quality of the coast.

Considering the very low SO_2 flux, the FLEXPART simulations in the CTRL run induce insignificant SO_2 concentration at an aircraft-sampled location (results not shown).

We performed CO simulations, as CO peaks were measured in one case out of two (Fig. 2c and f). For those simulations, we used the injection height of 27 m of the CTRL run and the flux from Deetz and Vogel (2017) of $1.1 \times 10^{-1} \text{ kg s}^{-1}$. For the flight on 10 July, FLEXPART sim-

ulates four CO peaks concomitant with the four NO_2 peaks, while no increase in CO was observed. On 14 July, three peaks of CO are simulated at the same time as the three measured ones but are underestimated. CO is always included as a gas emitted from flaring in the inventory for this specific vessel, while it does not seem to be the case each time. A discussion on the flaring combustion processes is presented below.

4.3 Combustion processes involved in oil flare

From a combustion point of view, flares generate non-premixed highly turbulent flames, characterised by high-frequency fluctuating flow fields. Flares can be air-assisted or steam-assisted in order to achieve a better efficiency. The

turbulence increases the mixing and affects the chemical reaction process. Very recent attempts to model such flames (Aboje et al., 2017; Damodara et al. 2017) showed that the influence of the composition of the natural gas and the validity of the chemical kinetic mechanism are of primary importance in predicting the emitted species. Moreover, the flame characteristics, such as temperature, height and length, have a strong impact on the dispersion of the plume together with wind speed and other meteorological variables (Rahnama and De Visscher, 2016). All these parameters need to be taken into account when considering air field campaigns.

McDaniel (1983) clearly demonstrated that CO emissions are strongly related to flare efficiency. When the efficiency is ca. 99.8 %, CO is below 10 ppmv (Fig. A-5 in McDaniel, 1983), while CO emissions may reach 440 ppmv when the efficiency drops to 64 % (Fig. A-8 in McDaniel, 1983). CO emissions can clearly be linked to the quality of the combustion, which is directly impacted by the turbulence. The observed difference between 10 and 14 July in terms of CO emissions mostly lies in the different wind conditions between these 2 d. First, the wind speed was lower on 14 July, which makes less O₂ available to burn with natural gas; second, it appears that the wind direction was not clearly established, as can be seen from the much more dispersed plume in Fig. 1b, resulting in incomplete combustion pockets favouring CO formation. However, a decrease in efficiency should also lead to lower temperatures and NO_x emissions, which is not observed here. The results of this campaign would require analysis in the light of computational fluid dynamic simulations, accounting for a realistic natural gas composition and its high-temperature chemistry, which are beyond the scope of this study.

5 Conclusions

This study was conducted in the framework of the DACCIIWA FP7 European project in July 2016 in southern West Africa. One target of the project was to measure emissions from oil rigs, which were not well estimated until then. With two flights planned in the vicinity of a FPSO platform, we reported the first flaring in situ measurements in this region. The aim of this work was to evaluate the capacity of the FLEXPART model to reproduce the NO₂ airborne measurements and to evaluate the inventory of Deetz and Vogel (2017) in the case of point sources of pollution such as oil platforms. The injection height of the plume was estimated by performing different calculations. According to several sensitivity runs, it appears that the emission flux given by Deetz and Vogel (2017) overestimates the concentrations, whereas a lower NO₂ flux roughly reproduces the measurements. Concerning the injection height, the sensitivity tests are not conclusive, showing the need for more and better-targeted measurements. An estimation of the pollutant distribution above or inside the MBL shows that the pollutants stay

mainly inside the MBL, limiting the transport to the coastline located 70 km downwind of the FPSO.

Sources of uncertainties are associated with the different calculations and hypotheses, but the work is mainly limited by the flight trajectories that are too far and too high from the platform. So it remains necessary to better quantify the emissions released during the flaring processes locally but also at wider scales. Generally speaking, this study suggests that for flights planned in the heart of a flaring plume, it should be possible to link the flaring observations obtained by satellites with the emissions deduced from the airborne measurements. If this relationship is possible, a general relationship between the emissions and the radiant heat could thus allow estimating the emissions of all flaring processes detected by satellite.

Data availability. The aircraft data used here can be accessed using the DACCIIWA database at <http://baobab.sedoo.fr/DACCIIWA/> (last access: October 2018; Brissebrat et al., 2017). A 2-year embargo period applies after the upload. External users can request the release of datasets before the end of the embargo period.

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Author contributions. HS was the mission scientist for the Falcon 20 aircraft and VC the principal investigator (PI) of the EUFAR2–APSOWA campaign. VC, SC and VB participated in the SPIRIT measurements aboard the aircraft, with remote assistance from CR for the SPIRIT maintenance. VB, GS and DS performed flight data analyses for their respective instruments. KD provided the updated version of his inventory for the period of the campaign and the VDI report. AD performed simulations using WRF-Chem to simulate the FPSO plume and gave advice for the modelling method. GK and VB performed the FLEXPART simulations. GD conducted the combustion part of the paper. VB wrote the paper, with contributions from all co-authors.

Competing interests. The authors declare that they have no conflict of interest.

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References

- Aboje, A. A., Hughes, K. J., Ingham, D. B., Ma, L., Williams, A., and Pourkashanian, M.: Numerical study of a wake-stabilized propane flame in a cross-flow of air, *Journal of Energy Institute*, 90, 145–158, <https://doi.org/10.1016/j.joei.2015.09.002>, 2017.
- Akingunola, A., Makar, P. A., Zhang, J., Darlington, A., Li, S.-M., Gordon, M., Moran, M. D., and Zheng, Q.: A chemical transport model study of plume-rise and particle size distribution for the Athabasca oil sands, *Atmos. Chem. Phys.*, 18, 8667–8688, <https://doi.org/10.5194/acp-18-8667-2018>, 2018.
- Anomohanran, O.: Determination of greenhouse gas emission resulting from gas flaring activities in Nigeria, *Energy Policy*, 45, 666–670, <https://doi.org/10.1016/j.enpol.2012.03.018>, 2012.
- Arya, S. P.: *Air Pollution Meteorology and Dispersion*, Oxford University Press, Department of Marine, Earth, and Atmospheric Sciences, North Carolina State University, 1999.
- Asuoha, A. N. and Osu Charles, I.: Seasonal variation of meteorological factors on air parameters and the impact of gas flaring on air quality of some cities in Niger Delta (Ibendo and its environs), *African Journal of Environmental Science and Technology*, 9, 218–227, <https://doi.org/10.5897/AJEST2015.1867>, 2015.
- Bahino, J., Yoboué, V., Galy-Lacaux, C., Adon, M., Akpo, A., Keita, S., Lioussé, C., Gardrat, E., Chiron, C., Ossouhou, M., Gnamien, S., and Djossou, J.: A pilot study of gaseous pollutants' measurement (NO_2 , SO_2 , NH_3 , HNO_3 and O_3) in Abidjan, Côte d'Ivoire: contribution to an overview of gaseous pollution in African cities, *Atmos. Chem. Phys.*, 18, 5173–5198, <https://doi.org/10.5194/acp-18-5173-2018>, 2018.
- Briggs, G. A.: A Plume Rise Model Compared with Observations, *Journal of the Air Pollution Control Association*, 15, 433–438, <https://doi.org/10.1080/00022470.1965.10468404>, 1965.
- Briggs, G. A.: Plume rise and buoyancy effects, atmospheric sciences and power production, in: DOE/TIC-27601 (DE84005177), edited by: Randerson, D., TN, Technical Information Center, U.S. Dept. of Energy, Oak Ridge, USA, 327–366, 1984.
- Brissebrat, G., Belmahfoud, N., Cloché, S., Ferré, H., Fleury, L., Mière, A., and Ramage, K.: The BAOBAB data portal and DACCIIWA database, 19th EGU General Assembly, EGU2017, 23–28 April, 2017, Vienna, Austria, p. 4685, available at: <http://baobab.sedoo.fr/DACCIIWA/> (last access: July 2018), 2017.
- Catoire, V., Robert, C., Chartier, M., Jacquet, P., Guimbaud, C., and Krysztofiak, G.: The SPIRIT airborne instrument: a three-channel infrared absorption spectrometer with quantum cascade lasers for in situ atmospheric trace-gas measurements, *Appl. Phys. B.*, 123, 244, <https://doi.org/10.1007/s00340-017-6820-x>, 2017.
- Damodara, V., Chen, D. H., Lou, H. H., Rasel, K. M. A., Richmond, P., Wang, A., and Li, X.: Reduced combustion mechanism for C_1 – C_4 hydrocarbons and its application in computational fluid dynamics flame modelling, *J. Air Waste Manage.*, 67, 599–612, <https://doi.org/10.1080/10962247.2016.1268546>, 2017.
- Dee, D. P., Uppala, S. M., Simmons, A. J., Berrisford, P., Poli, P., Kobayashi, S., Andrae, U., Balmaseda, M. A., Balsamo, G., Bauer, P., Bechtold, P., Beljaars, A. C. M., van de Berg, L., Bidlot, J., Bormann, N., Delsol, C., Dragani, R., Fuentes, M., Geer, A. J., Haimberger, L., Healy, S. B., Hersbach, H., Hólm, E. V., Isaksen, I., Kållberg, P., Köhler, M., Matricardi, M., McNally, A. P., Monge-Sanz, B. M., Morcrette, J.-J., Park, B.-K., Peubey, C., de Rosnay, P., Tavolato, C., Thépaut, J.-N., and Vitart, F.: The ERA-Interim reanalysis: configuration and performance of the data assimilation system, *Q. J. Roy. Meteor. Soc.*, 137, 553–597, <https://doi.org/10.1002/qj.828>, 2011.
- Deetz, K. and Vogel, B.: Development of a new gas-flaring emission dataset for southern West Africa, *Geosci. Model Dev.*, 10, 1607–1620, <https://doi.org/10.5194/gmd-10-1607-2017>, 2017.
- Elvidge, C. D., Zhizhin, M., Hsu, F.-C., and Baugh, K. E.: VIIRS Nightfire: Satellite Pyrometry at Night, *Remote Sensing*, 5, 4423–4449, <https://doi.org/10.3390/rs5094423>, 2013.
- Elvidge, C. D., Zhizhin, M., Baugh, K., Hsu, F.-C., and Ghosh, T.: Methods for Global Survey of Natural Gas Flaring from Visible Infrared Imaging Radiometer Suite Data, *Energies*, 9, 14, <https://doi.org/10.3390/en9010014>, 2015.
- Fawole, O. G., Cai, X., Levine, J. G., Pinker, R. T., and MacKenzie, A. R.: Detection of a gas flaring signature in the AERONET optical properties of aerosols at a tropical station in West Africa, *J. Geophys. Res.-Atmos.*, 121, 2016JD025584, <https://doi.org/10.1002/2016JD025584>, 2016a.
- Fawole, O. G., Cai, X.-M., and MacKenzie, A. R.: Gas flaring and resultant air pollution: A review focusing on black carbon, *Environ. Pollut.*, 216, 182–197, <https://doi.org/10.1016/j.envpol.2016.05.075>, 2016b.
- Fiebig, M.: *Das troposphärische Aerosol in mittleren Breiten – Mikrophysik, Optik und Klimaantrieb am Beispiel der Feldstudie LACE 98*, PhD thesis, Ludwig-Maximilians-Universität, Institut für Physik der Atmosphäre, DLR, Oberpfaffenhofen, 2001.
- Fiebig, M., Stein, C., Schroder, F., Feldpausch, P., and Petzold, A.: Inversion of data containing information on the aerosol particle size distribution using multiple instruments, *J. Aerosol Sci.*, 36, 1353, <https://doi.org/10.1016/j.jaerosci.2005.01.004>, 2005.
- Flamant, C., Knippertz, P., Fink, A. H., Akpo, A., Brooks, B., Chiu, C. J., Coe, H., Danuor, S., Evans, M., Jegede, O., Kalthoff, N., Konaré, A., Lioussé, C., Lohou, F., Mari, C., Schlager, H., Schwarzenboeck, A., Adler, B., Amekudzi, L., Aryee, J., Ayoola, M., Batenburg, A. M., Bessardon, G., Borrmann, S.,

- Brito, J., Bower, K., Burnet, F., Catoire, V., Colomb, A., Denjean, C., Fosu-Amankwah, K., Hill, P. G., Lee, J., Lothon, M., Maranan, M., Marsham, J., Meynadier, R., Ngamini, J., Rosenberg, P., Sauer, D., Smith, V., Stratmann, G., Taylor, J. W., Voigt, C., and Yoboué, V.: The Dynamics-Aerosol-Chemistry-Cloud Interactions in West Africa field campaign: Overview and research highlights, *B. Am. Meteorol. Soc.*, 99, 83–104, <https://doi.org/10.1175/BAMS-D-16-0256.1>, 2017.
- Guigard, S. E., Kindzierski, W. B., and Harper, N.: Heat Radiation from Flares. Report prepared for Science and Technology Branch, Alberta Environment, ISBN 0-7785-1188-X, Edmonton, Alberta, 2000.
- Hassan, A. and Kouhy, R.: Gas flaring in Nigeria: Analysis of changes in its consequent carbon emission and reporting, *Accounting Forum*, 37, 124–134, <https://doi.org/10.1016/j.accfor.2013.04.004>, 2013.
- Ismail, O. S. and Umukoro, G. E.: Modelling combustion reactions for gas flaring and its resulting emissions, *Journal of King Saud University – Eng. Sci.*, 28, 130–140, <https://doi.org/10.1016/j.jksues.2014.02.003>, 2016.
- Knippertz, P., Coe, H., Chiu, J. C., Evans, M. J., Fink, A. H., Kalhoff, N., Lioussé, C., Mari, C., Allan, R. P., Brooks, B., Danour, S., Flamant, C., Jegede, O. O., Lohou, F., and Marsham, J. H.: The DACCWA Project: Dynamics–Aerosol–Chemistry–Cloud Interactions in West Africa, *B. Am. Meteorol. Soc.*, 96, 1451–1460, <https://doi.org/10.1175/BAMS-D-14-00108.1>, 2015.
- Law, K. S., Roiger, A., Thomas, J. L., Marelle, L., Raut, J. C., Dalsøren, S., Fuglestad, J., Tuccella, P., Weinzierl, B., and Schlager, H.: Local Arctic air pollution: Sources and impacts, *Ambio*, 46, 453–463, <https://doi.org/10.1007/s13280-017-0962-2>, 2017.
- McDaniel, M.: Flare efficiency study, available at: http://www.epa.gov/ttn/chief/ap42/ch13/related/ref_01c13s05_jan1995.pdf (last access: July 2018), 1983.
- Nwankwo, C. N. and Ogagarue, D. O.: Effects of gas flaring on surface and ground waters in Delta State Nigeria, *Journal of Geology and Mining Research*, 3, 131–136, 2011.
- Nwaugo, V. O., Onyeagba, R. A., and Nwahukwu, N. C.: Effect of gas flaring on soil microbial spectrum in parts of Niger Delta area of southern Nigeria, *Afr. J. Biotechnol.*, 5, 1824–1826, 2006.
- Osuji, L. C. and Avwiri, G. O.: Flared Gas and Other Pollutants Associated with Air quality in industrial Areas of Nigeria: An Overview, *Chem. Biodivers.*, 2, 1277–1289, 2005.
- Rahnama, K. and De Visscher, A.: Simplified Flare Combustion Model for Flare Plume Rise Calculations, *Can. J. Chem. Eng.*, 94, 1249–1261, 2016.
- Rothman, L. S., Gordon, I. E., Babikov, Y., Barbe, A., Chris Benner, D., Bernath, P. F., Birk, M., Bizzocchi, L., Boudon, V., Brown, L. R., Campargue, A., Chance, K., Cohen, E. A., Coudert, L. H., Devi, V. M., Drouin, B. J., Fayt, A., Flaud, J.-M., Gamache, R. R., Harrison, J. J., Hartmann, J.-M., Hill, C., Hodges, J. T., Jacquemart, D., Jolly, A., Lamouroux, J., Le Roy, R. J., Li, G., Long, D. A., Lyulin, O. M., Mackie, C. J., Massie, S. T., Mikhailenko, S., Müller, H. S. P., Naumenko, O. V., Nikitin, A. V., Orphal, J., Perevalov, V., Perrin, A., Polovtseva, E. R., Richard, C., Smith, M. A. H., Starikova, E., Sung, K., Tashkun, S., Tennyson, J., Toon, G. C., Tyuterev, V. G., and Wagner, G.: The HITRAN2012 molecular spectroscopic database, *J. Quant. Spectrosc. Ra.*, 130, 4–50, <https://doi.org/10.1016/j.jqsrt.2013.07.002>, 2013.
- Schröder, F. and Ström, J.: Aircraft measurements of sub-micrometer aerosol particles (> 7 nm) in the midlatitude free troposphere and tropopause region, *Atmos. Res.*, 44, 333, [https://doi.org/10.1016/S0169-8095\(96\)00034-8](https://doi.org/10.1016/S0169-8095(96)00034-8), 1997.
- Shimamura, Y.: FPSO/FSO: State of the art, *J. Mar. Sci. Technol.*, 7, 59–70, <https://doi.org/10.1007/s007730200013>, 2002.
- Sonibare, J. A. and Akeredolu, F. A.: A theoretical prediction of non-methane gaseous emissions from natural gas combustion, *Energy Policy*, 32, 1653–1665, <https://doi.org/10.1016/j.enpol.2004.02.008>, 2004.
- Stohl, A., Forster, C., Frank, A., Seibert, P., and Wotawa, G.: Technical note: The Lagrangian particle dispersion model FLEXPART version 6.2, *Atmos. Chem. Phys.*, 5, 2461–2474, <https://doi.org/10.5194/acp-5-2461-2005>, 2005.
- Tuccella, P., Thomas, J. L., Law, K. S., Raut, J.-C., Marelle, L., Roiger, A., Weinzierl, B., Denier van der Gon, H.A.C., Schlager, H., and Onishi, T.: Air pollution impacts due to petroleum extraction in the Norwegian Sea during the ACCESS aircraft campaign, *Elem. Sci. Anth.*, 5, 25, <https://doi.org/10.1525/elementa.124>, 2017.
- VDI 3782: Dispersion of Air Pollutants in the Atmosphere, Determination of Plume rise, Verein Deutscher Ingenieure, VDI-Richtlinien 3782 Part 3, 1985.
- Villasenor, R., Magdaleno, M., Quintanar, A., Gallardo, J. C., López, M., Jurado, R., Miranda, A., Aguilar, M., Melgarejo, L. A., Palmerín, E., Vallejo, C. J., and Barchet, W. R.: An air quality emission inventory of offshore operations for the exploration and production of petroleum by the Mexican oil industry, *Atmos. Environ.*, 37, 3713–3729, [https://doi.org/10.1016/S1352-2310\(03\)00445-X](https://doi.org/10.1016/S1352-2310(03)00445-X), 2003.
- von der Weiden, S.-L., Drewnick, F., and Borrmann, S.: Particle Loss Calculator – a new software tool for the assessment of the performance of aerosol inlet systems, *Atmos. Meas. Tech.*, 2, 479–494, <https://doi.org/10.5194/amt-2-479-2009>, 2009.
- Zhang, Y., Gautam, R., Zavala-Araiza, D., Jacob, D. J., Zhang, R., Zhu, L., Sheng J.-X., and Scarpelli, T.: Satellite-observed changes in Mexico's offshore gas flaring activity linked to oil/gas regulations, *Geophys. Res. Lett.*, 46, 1879–1888, <https://doi.org/10.1029/2018GL081145>, 2019.